



***Earth's surface movements in relation to Parkfield  
2004 earthquake: Interpretation  
of permanent GPS observations***

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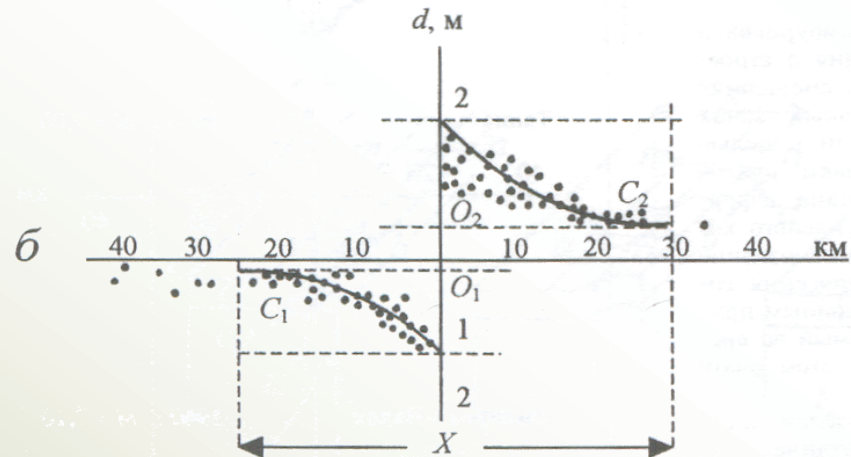
**IAG Scientific Assembly, Potsdam, September 1-6, 2013**

# Agelong history of elastic rebound model

- San Francisco earthquake, April 18, 1906,  $M=7.9$ ,
- Reid, H.F., The Mechanics of the Earthquake, The California Earthquake of April 18, 1906, Report of the State Investigation Commission, Vol.2, Carnegie Institution of Washington, Washington, D.C. **1910**
- Kanamori H. Mode of strain release associated with major earthquake in Japan / In: Donath F.A. (editor) Annual Review of Earth and Planetary Sciences, Annual Reviews, Palo Alto, Calif. – 1972.- 1.- 213
- Pevnev A. K. , “Earthquake prediction: geodetic aspects of the problem,” Izvestiya, Physics of the Solid Earth, vol. 12, pp. 88–98, 1988.



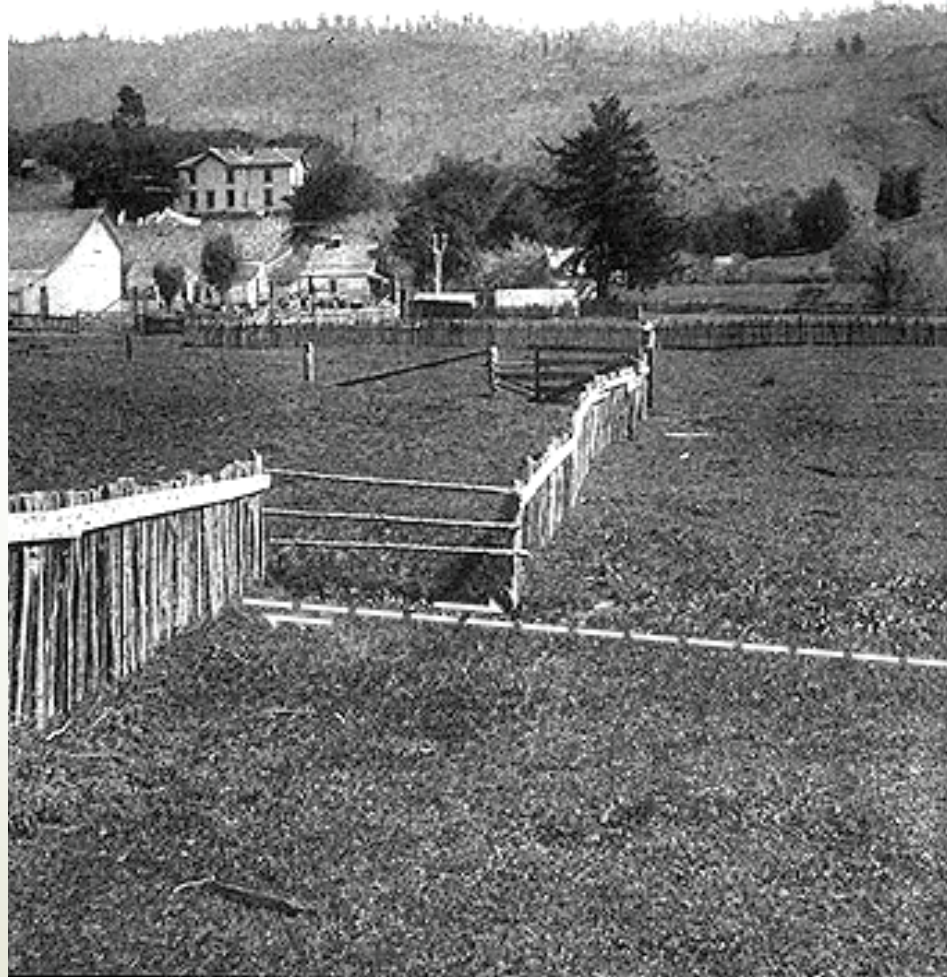
# Classical model example of Tango earthquake (Japan, 1927, $M=7.5$ )



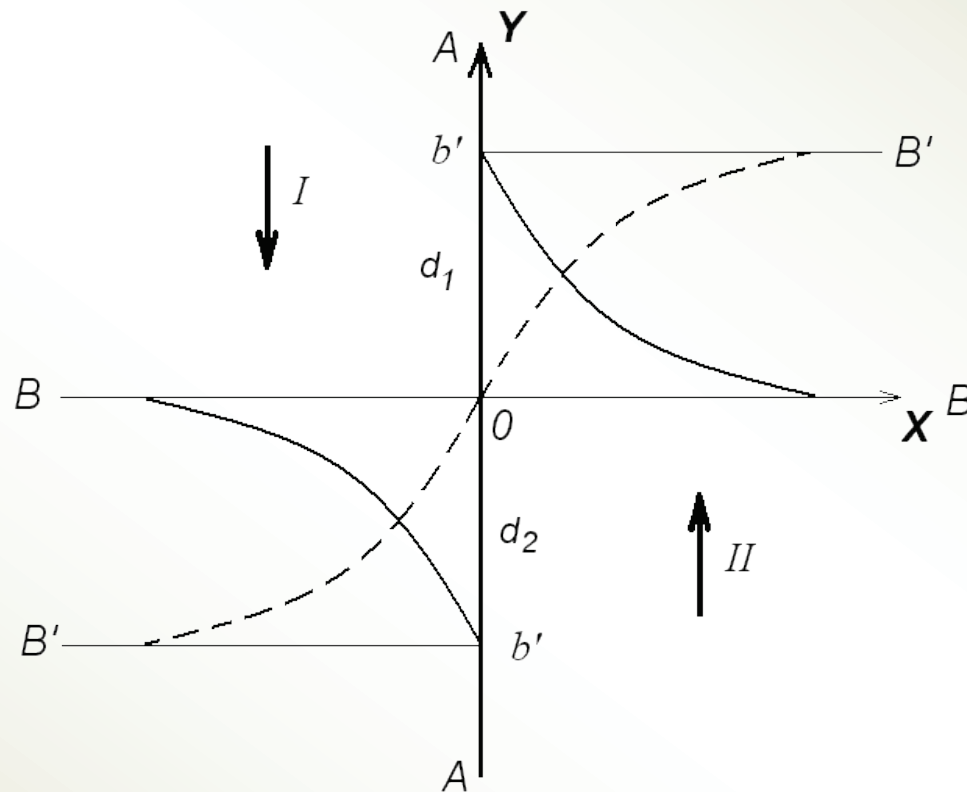




# The fence on the San Andreas (by G.K. Gilbert from Steinbrugge Collection of the UC Berkeley Earthquake Engineering Research Center)

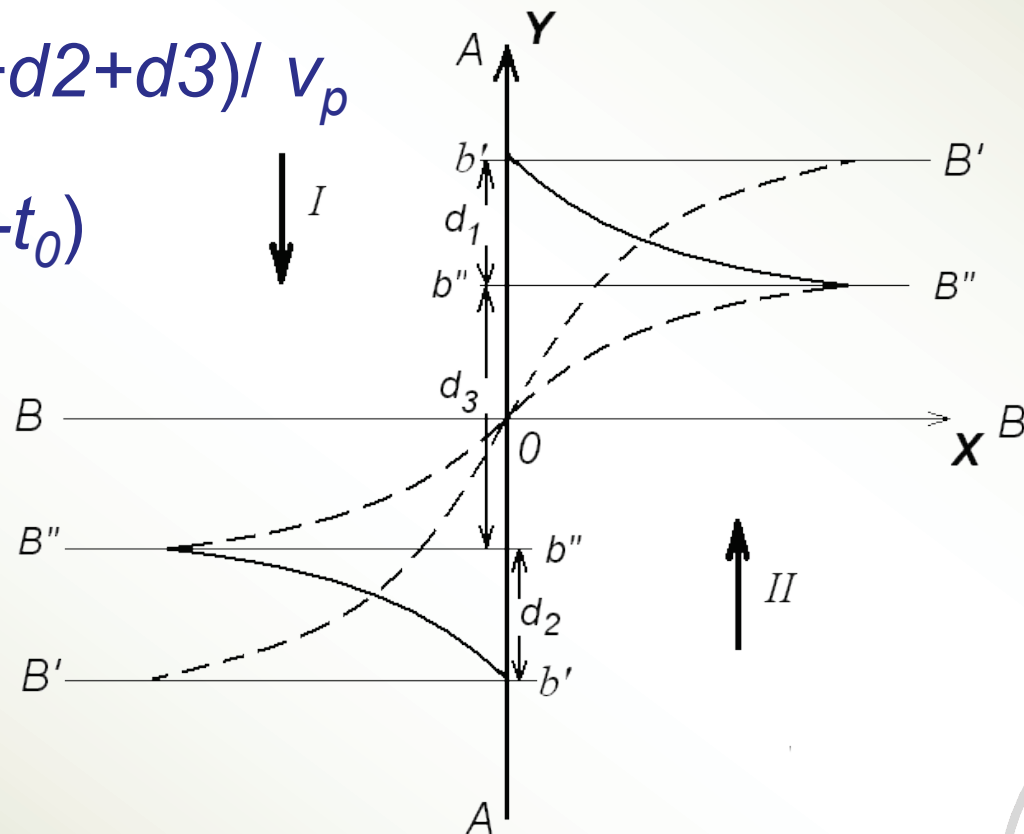


# Elastic rebound (case 2: first epoch a day before the earthquake)



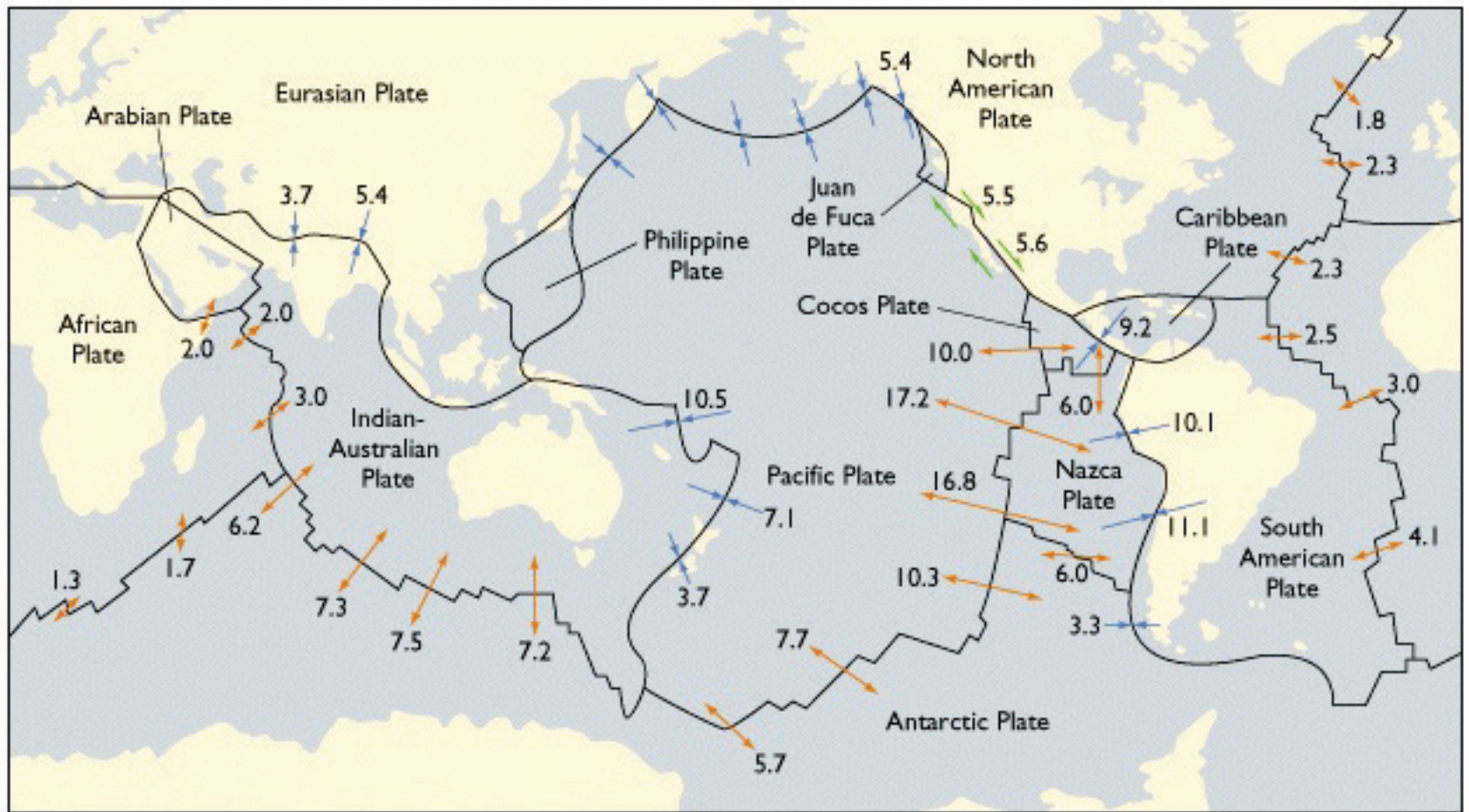
# Elastic rebound (common case 3: first epoch at the deformed surface)

- $T_p = (d_1 + d_2 + d_3) / v_p$
- $v_p = d_3 / (t_i - t_0)$

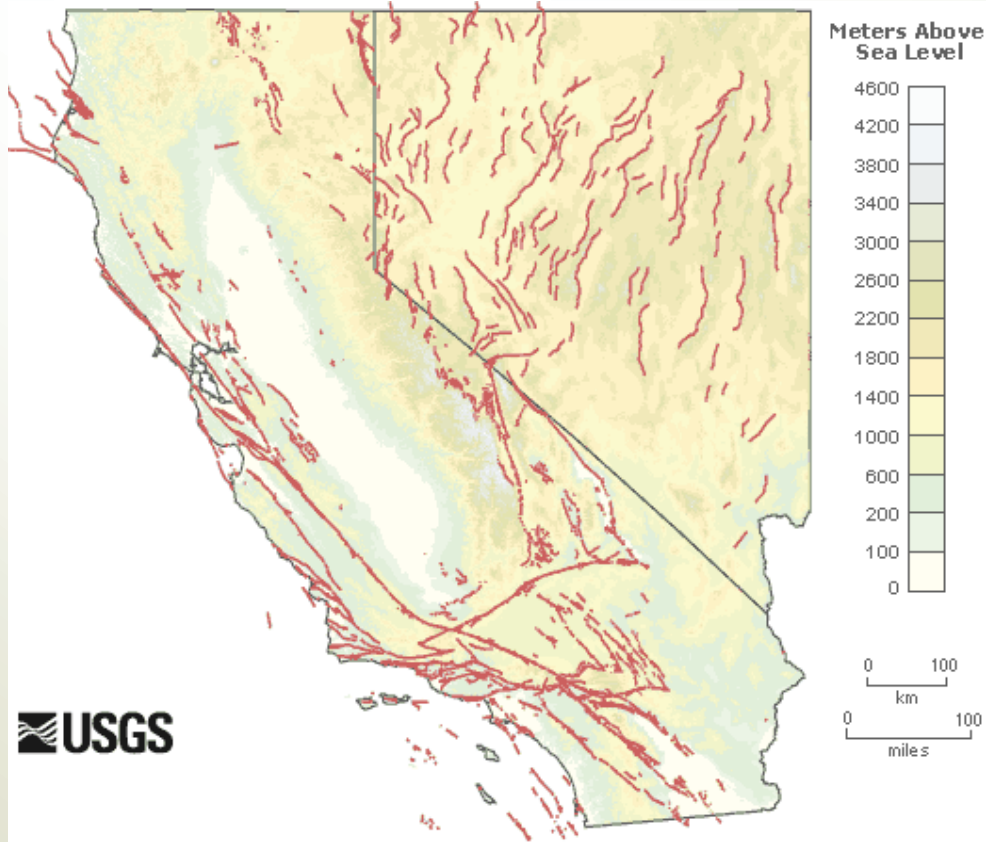




The lithospheric plate boundaries are of three types:  
Divergent, Convergent, Transform.  
The model more usable to the last one



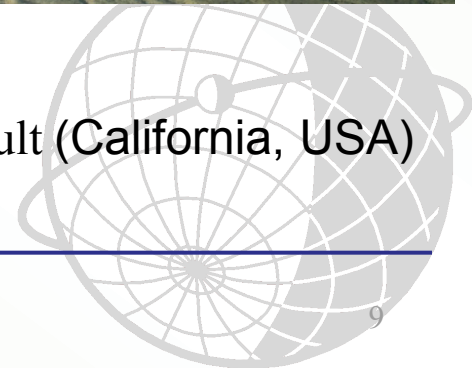
# Investigated region



Tectonic fault map (California, USA)



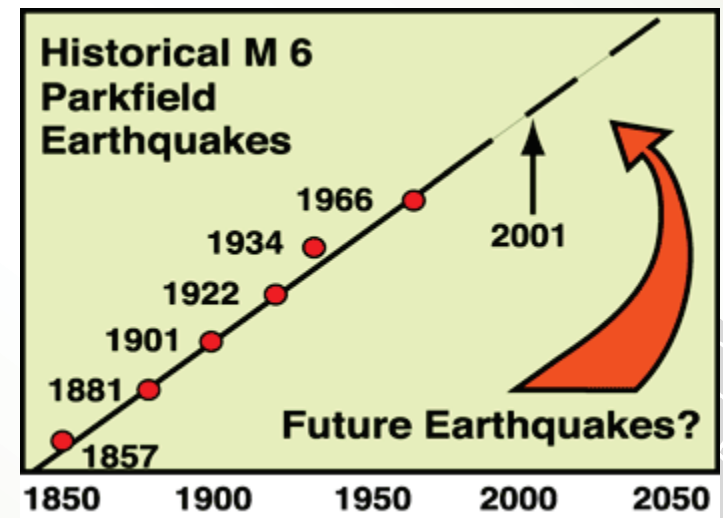
San Andreas fault (California, USA)





# Parkfield prediction experiment. California Integrated Seismic Network.

- Long term prediction for the strong earthquake moment was 1985-1993
- Prediction was based on about regular earthquake repeatability
- Place was predicted accurately





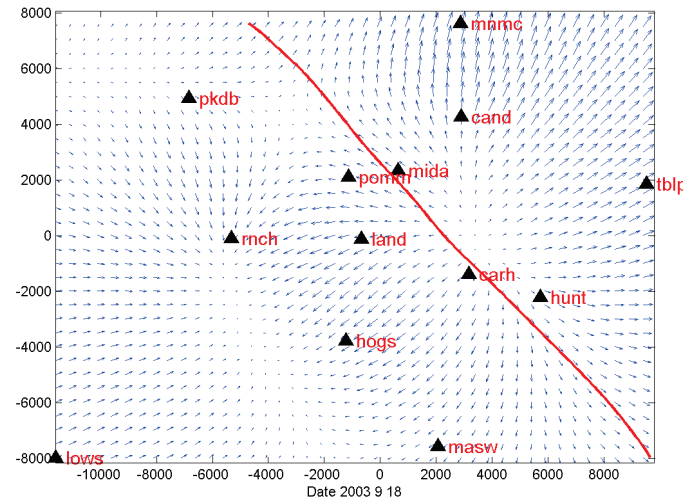
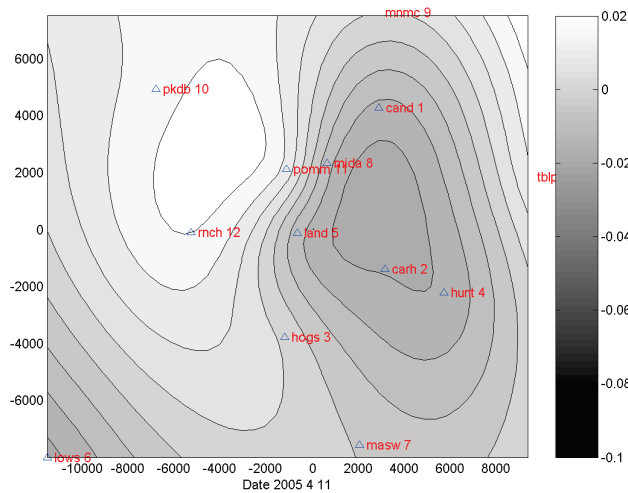
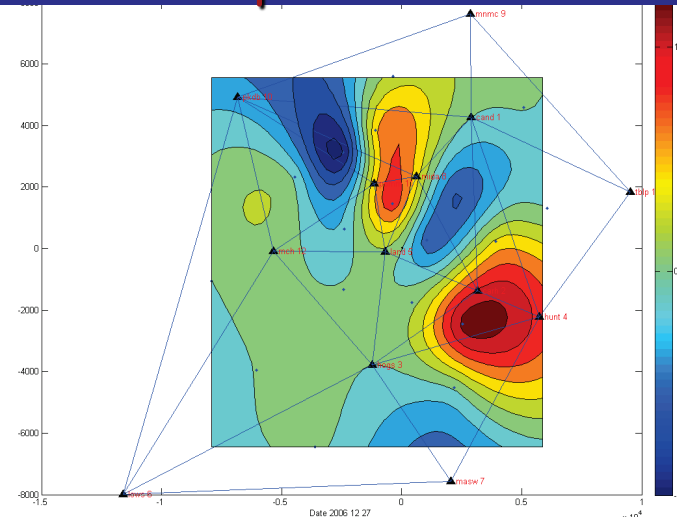
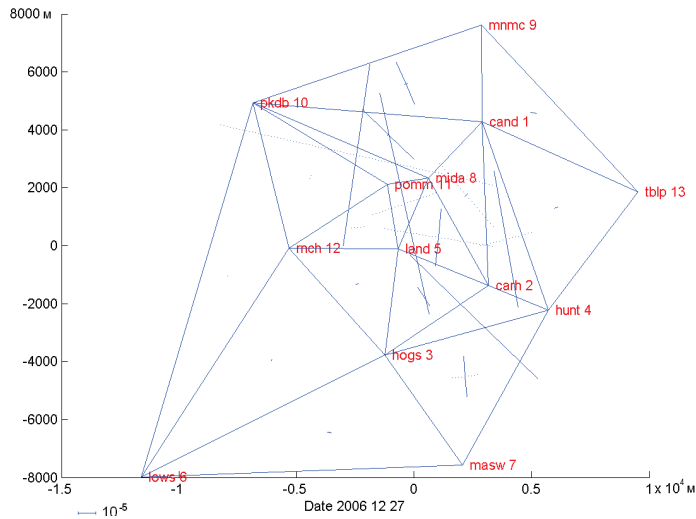
# GPS network specification

Number of points:	13
Number of selected baselines:	30
Baseline lengths:	1—14 km
Organization:	PBO (Plate Boundary Observatory)
Measurement mode:	permanent

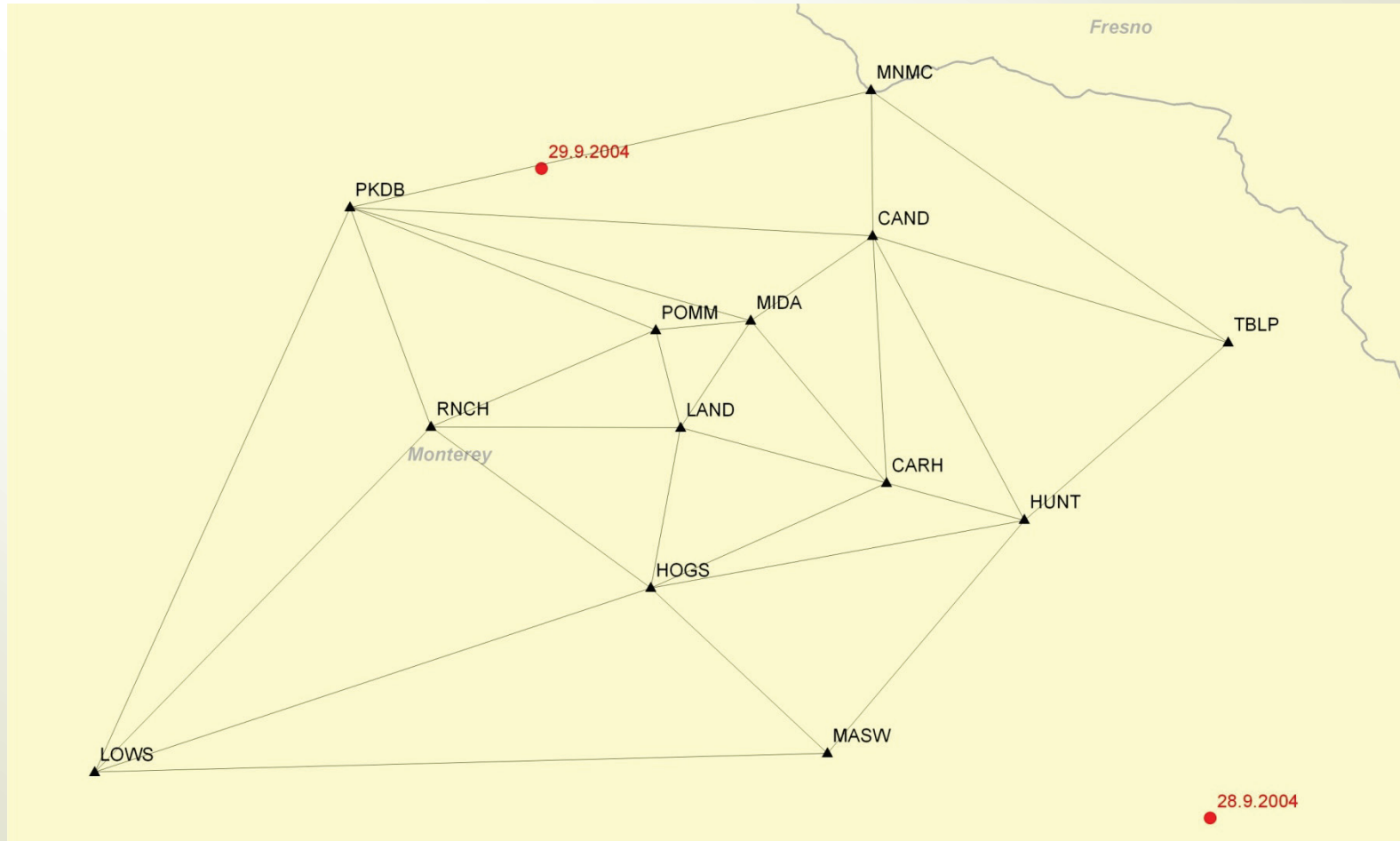


# Received deformation:

- 1) Main deformation axes
- 2) Dilatation
- 3) Vertical displacements
- 4) Horizontal gradients of vertical displacements



# PBO GPS network (Parkfield experiment tool)



Main Parkfield earthquake epicenters, September 28-29 2004 (M 6.0 and 5.1)

# Measurement results received

- SOPAC archive data was used for baseline processing
- Used date interval: 01.01.2002—27.12.2006
- Number of processed epochs: 365
- Interval between epochs: 5d
- Used software: Topcon Tools v. 7.1



# Earth crust movement and deformation analysis

- Initial data:
  - Baseline vectors
  - Covariance matrixes
- Data processing (the special software was made and used)
  - Temporal difference baseline components computation
  - Covariance matrix determination
  - Especial adjustment for determination of point displacements



# Earth crust movement and deformation analysis

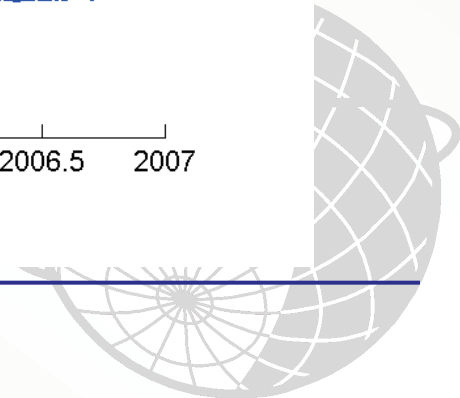
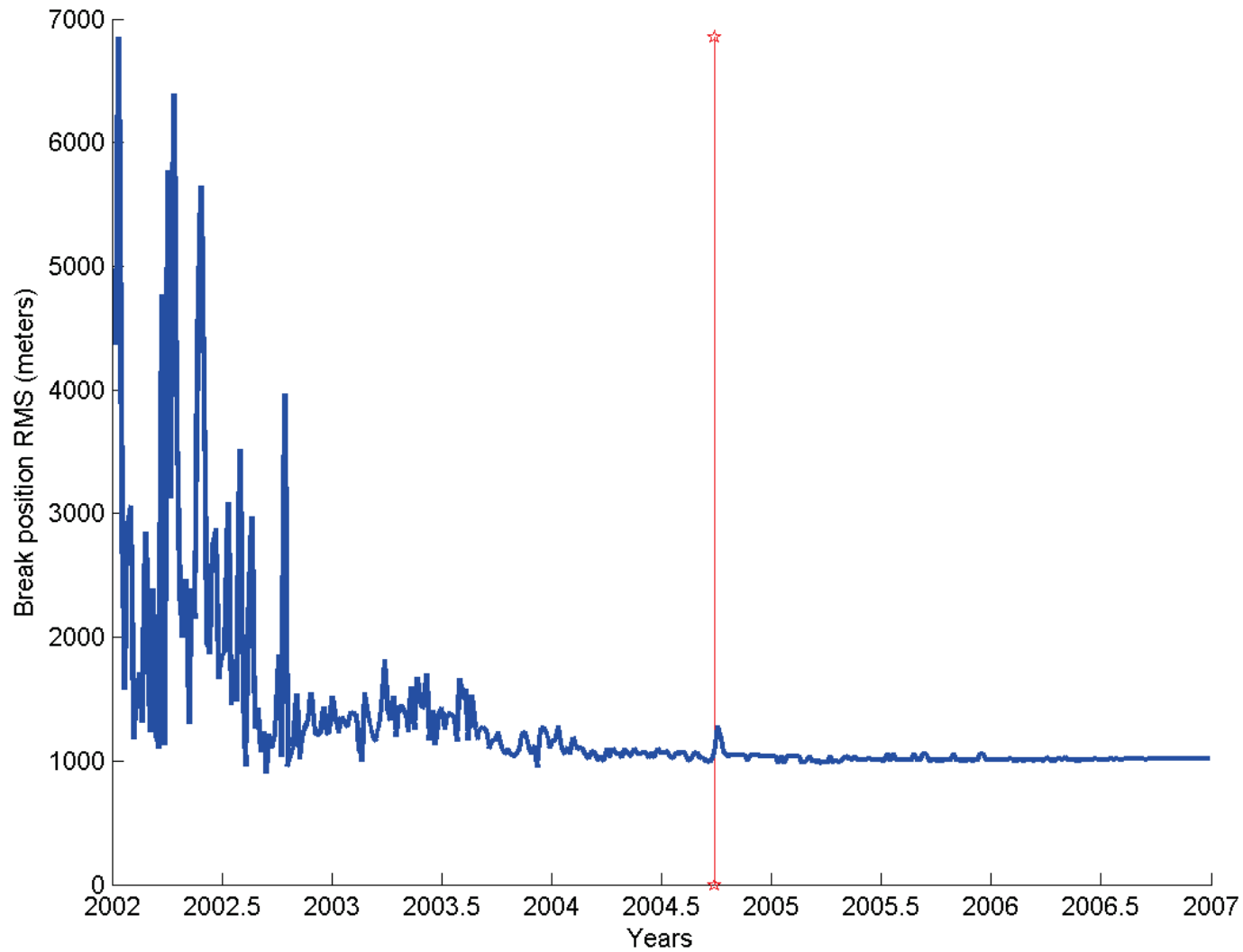
- Deformation analysis
  - Transformation to a local reference system (NEU)
  - Deformation component computation
  - Accuracy estimation
- Graphic and animation representation



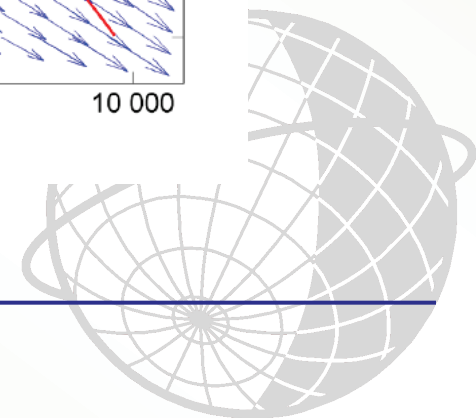
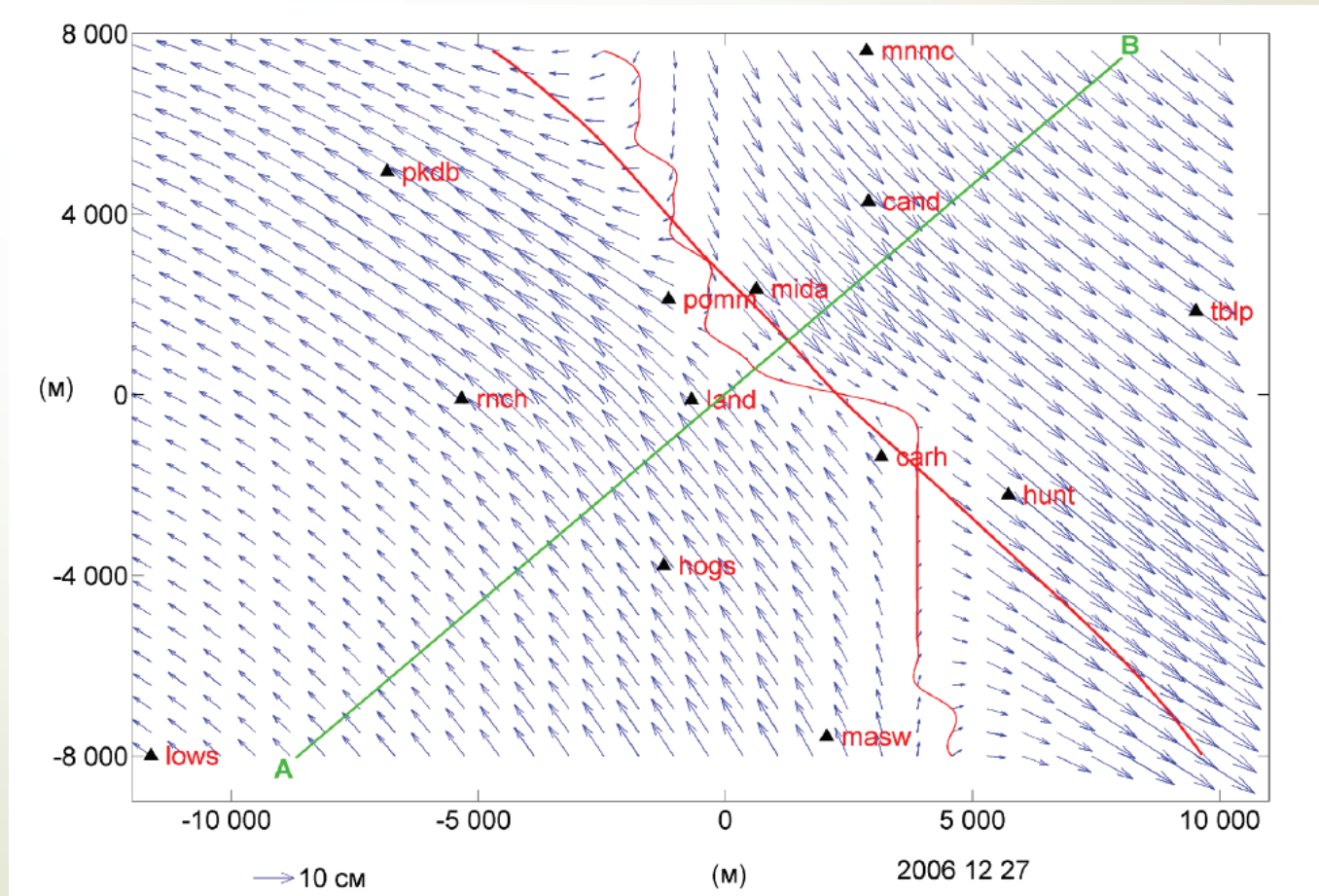




# Rupture position RMS

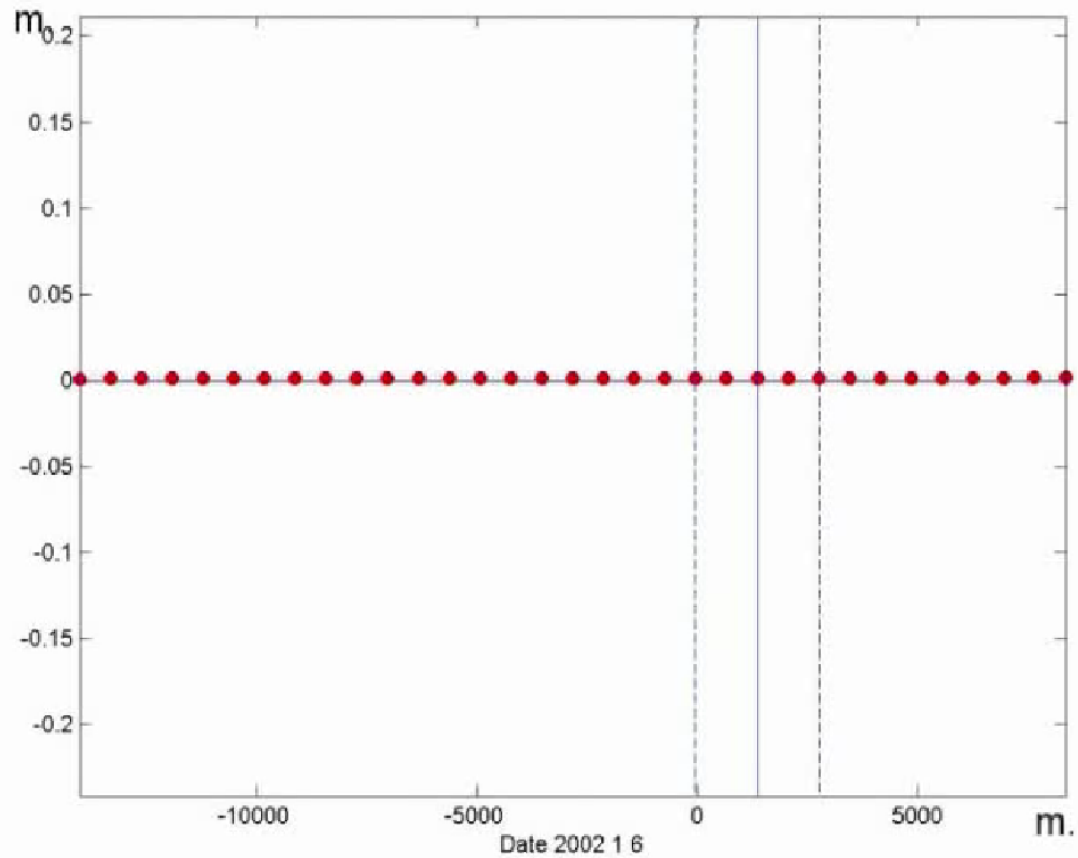


# Elastic rebound test cross-section



# Elastic rebound animation

Fault zone



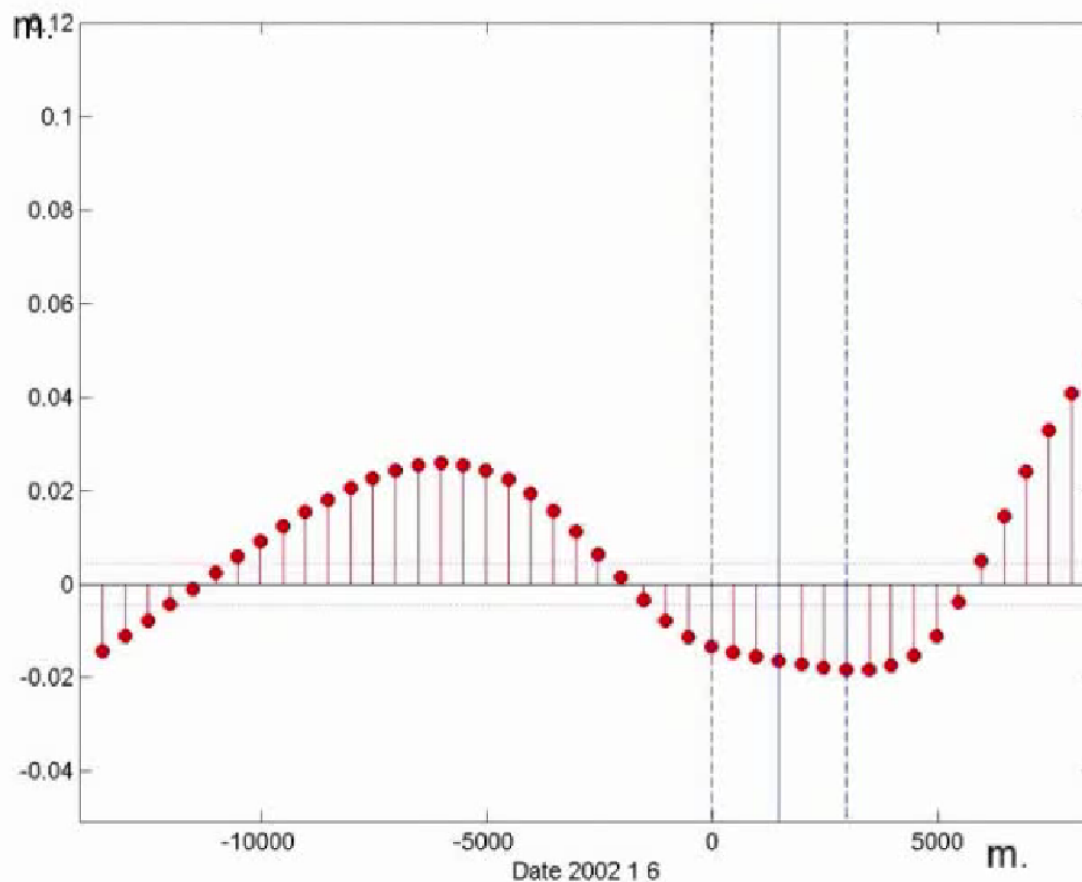
# Computation of strain accumulation start moment

- $t = 2.74$  yr
- $dN_i - dN_j = 42.5$  cm
- $dN_i' - dN_j' = 3.5$  cm
- **$T = (dN_i - dN_j)t / (dN_i' - dN_j') = 33 \pm 7$  yr**
- Elastic strain accumulation has begun at  $1972 \pm 7$  just after the previous earthquake of 1966



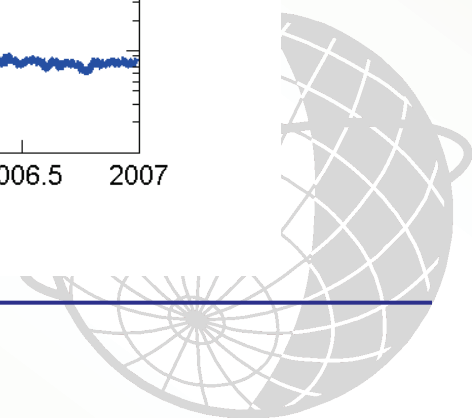
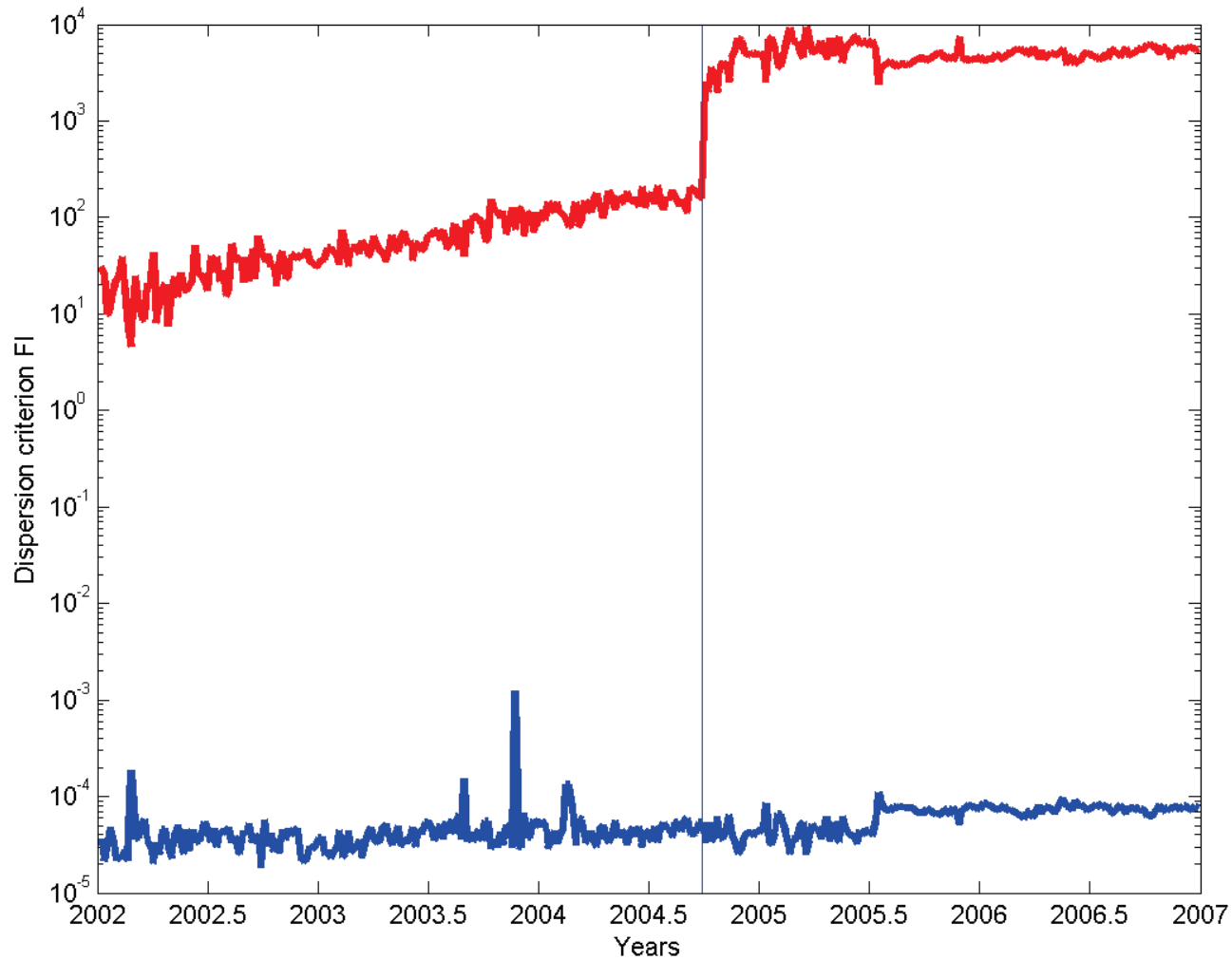
# Vertical displacements across the fault

Fault zone

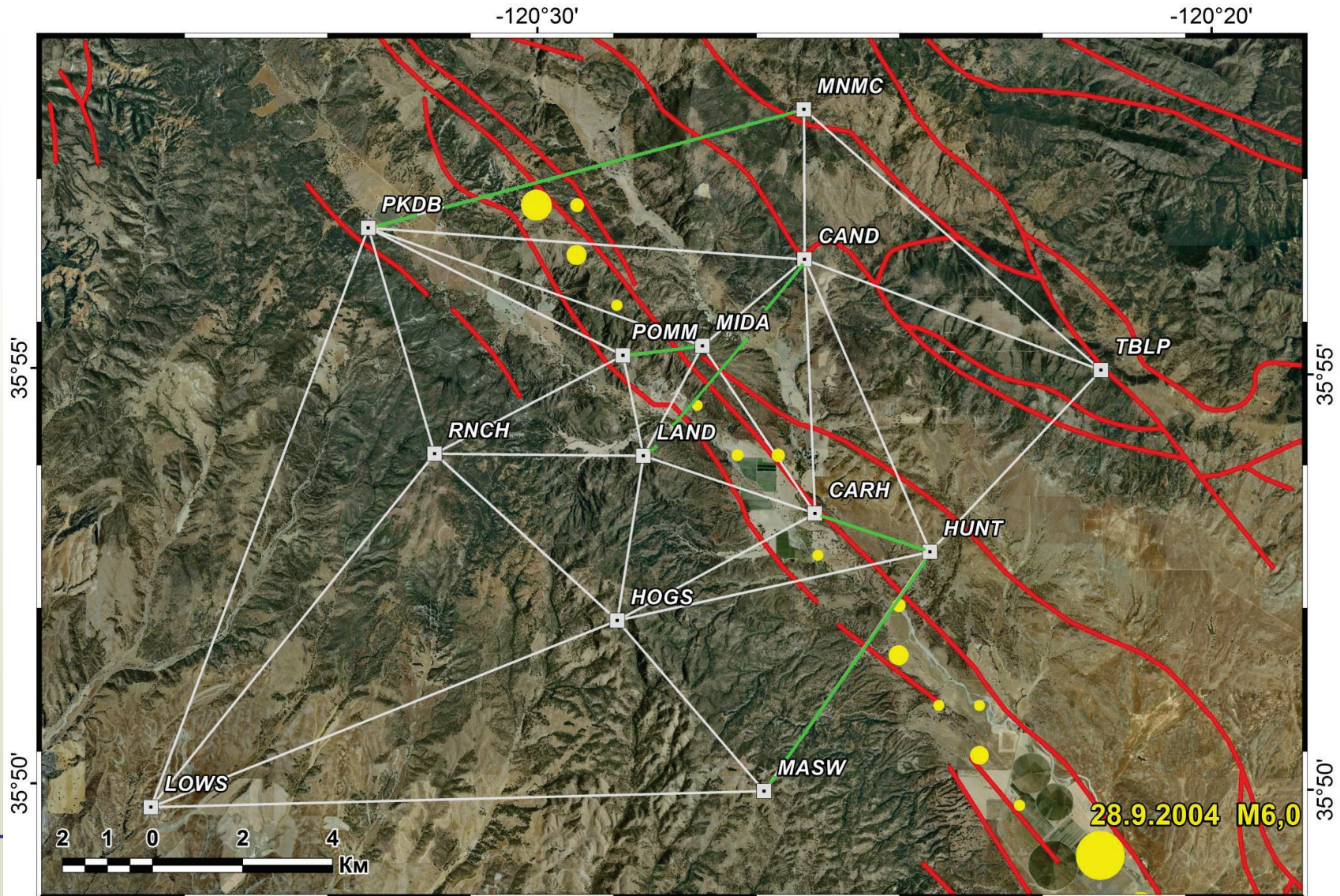




# Global deformation test

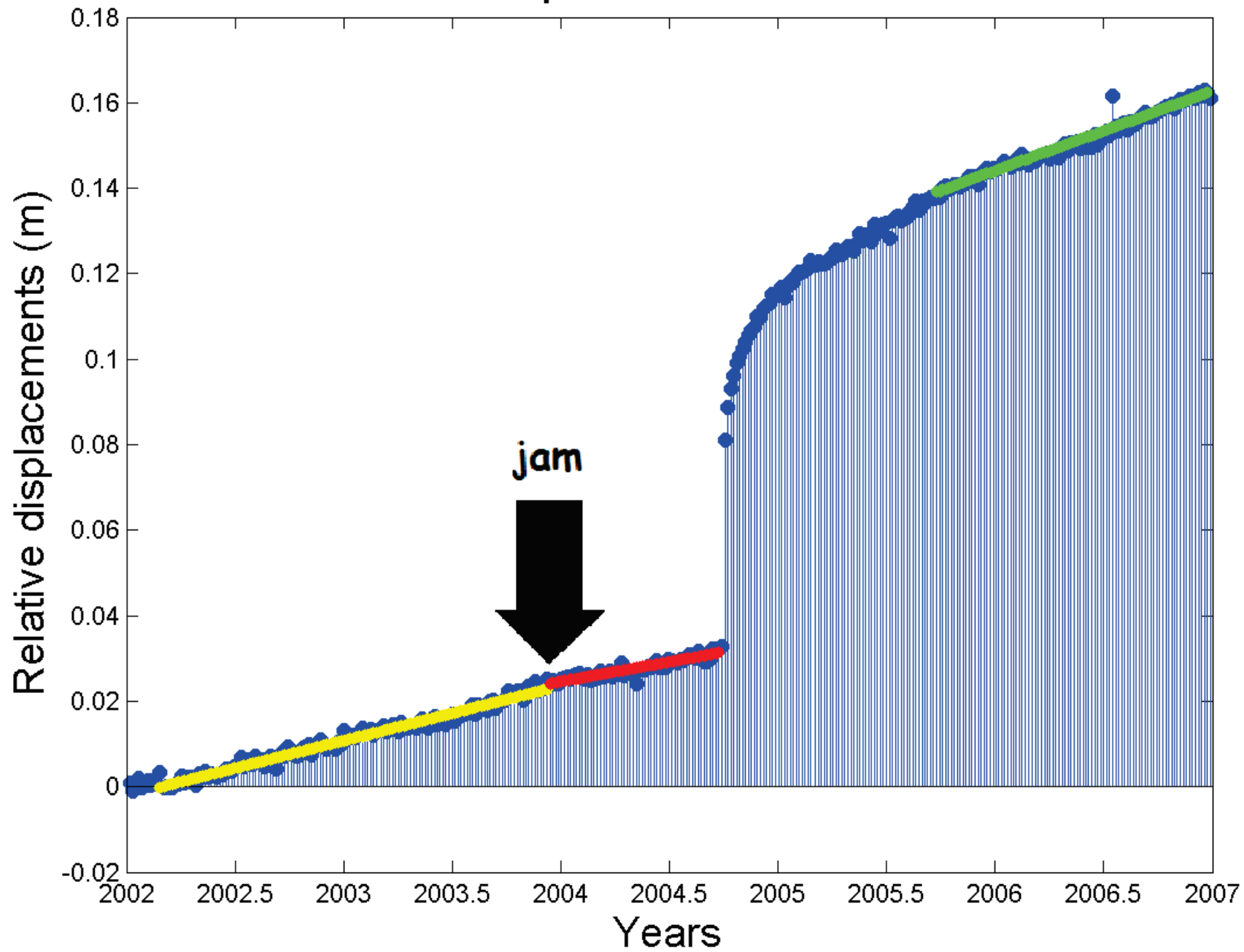


# Creep velocity comparison moving to epicentre

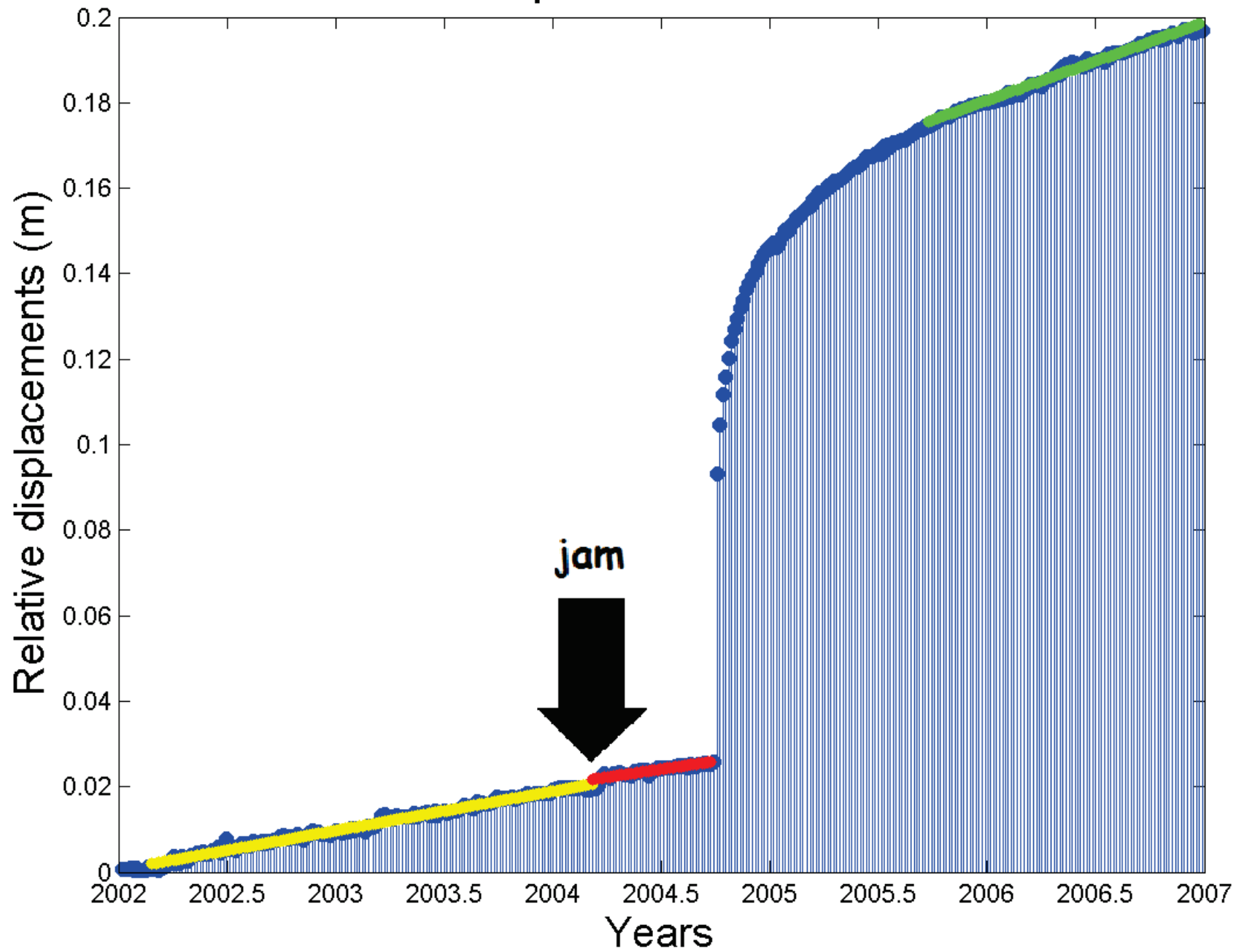




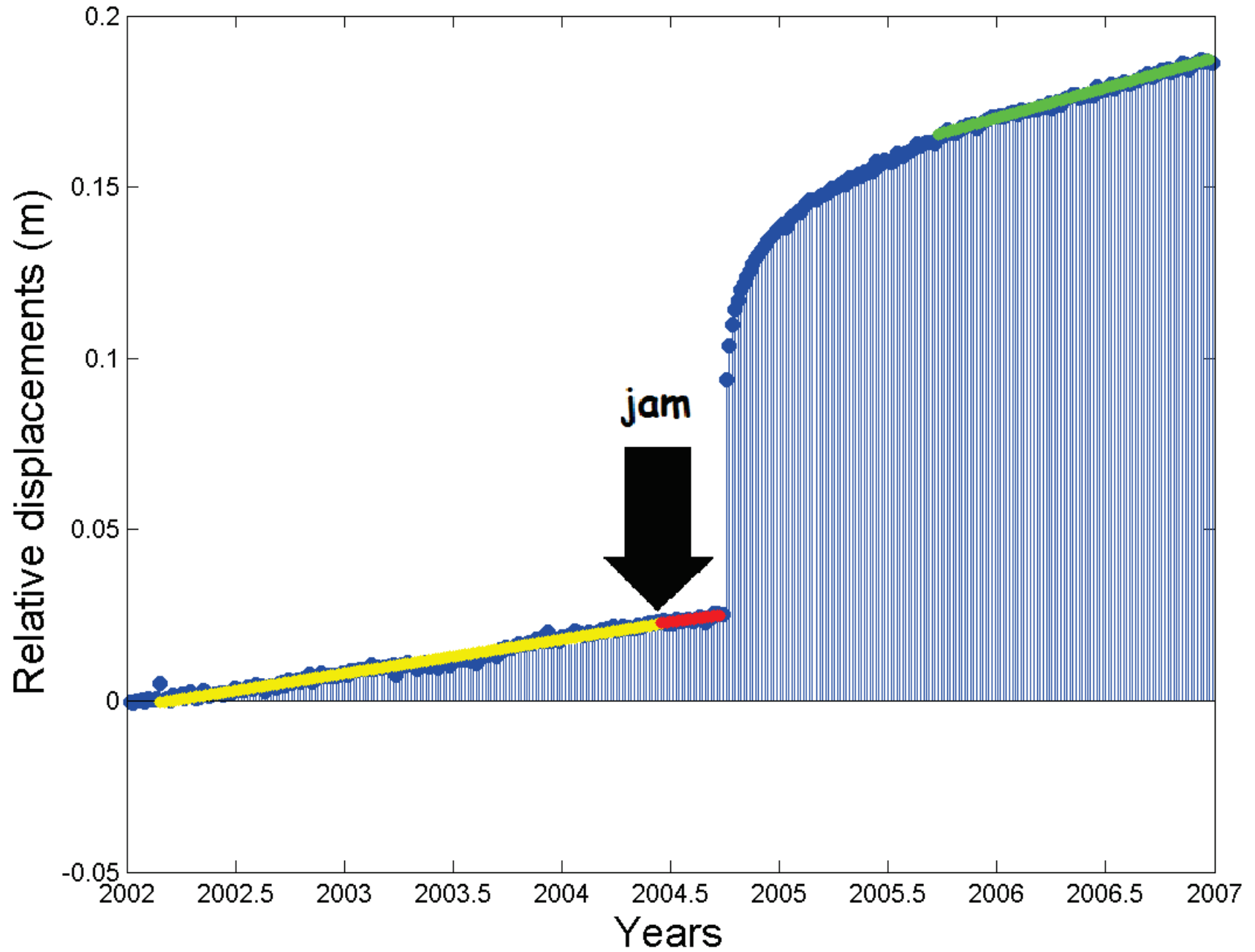
# pkdb-mnmc



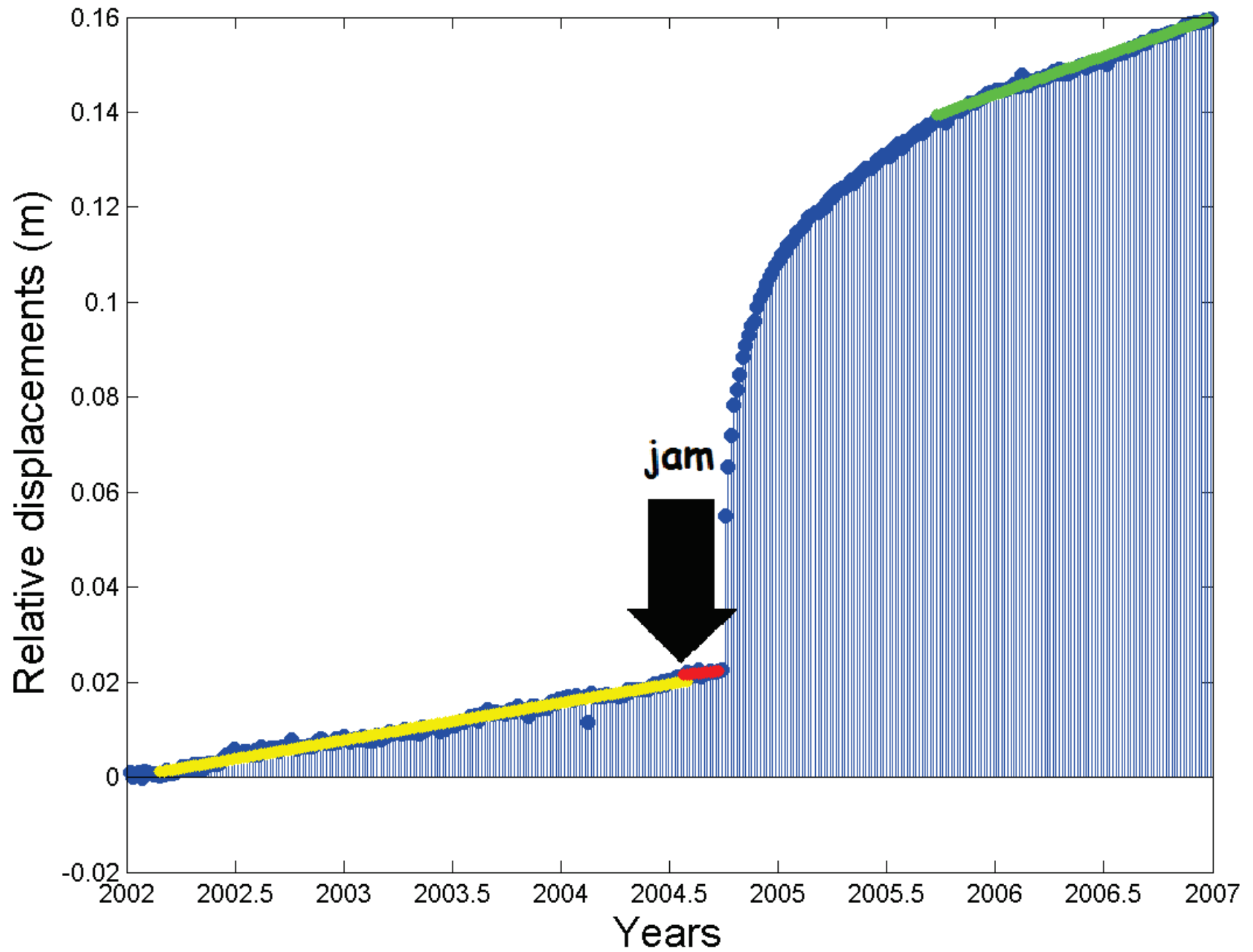
# pomm-mida



# land-cand

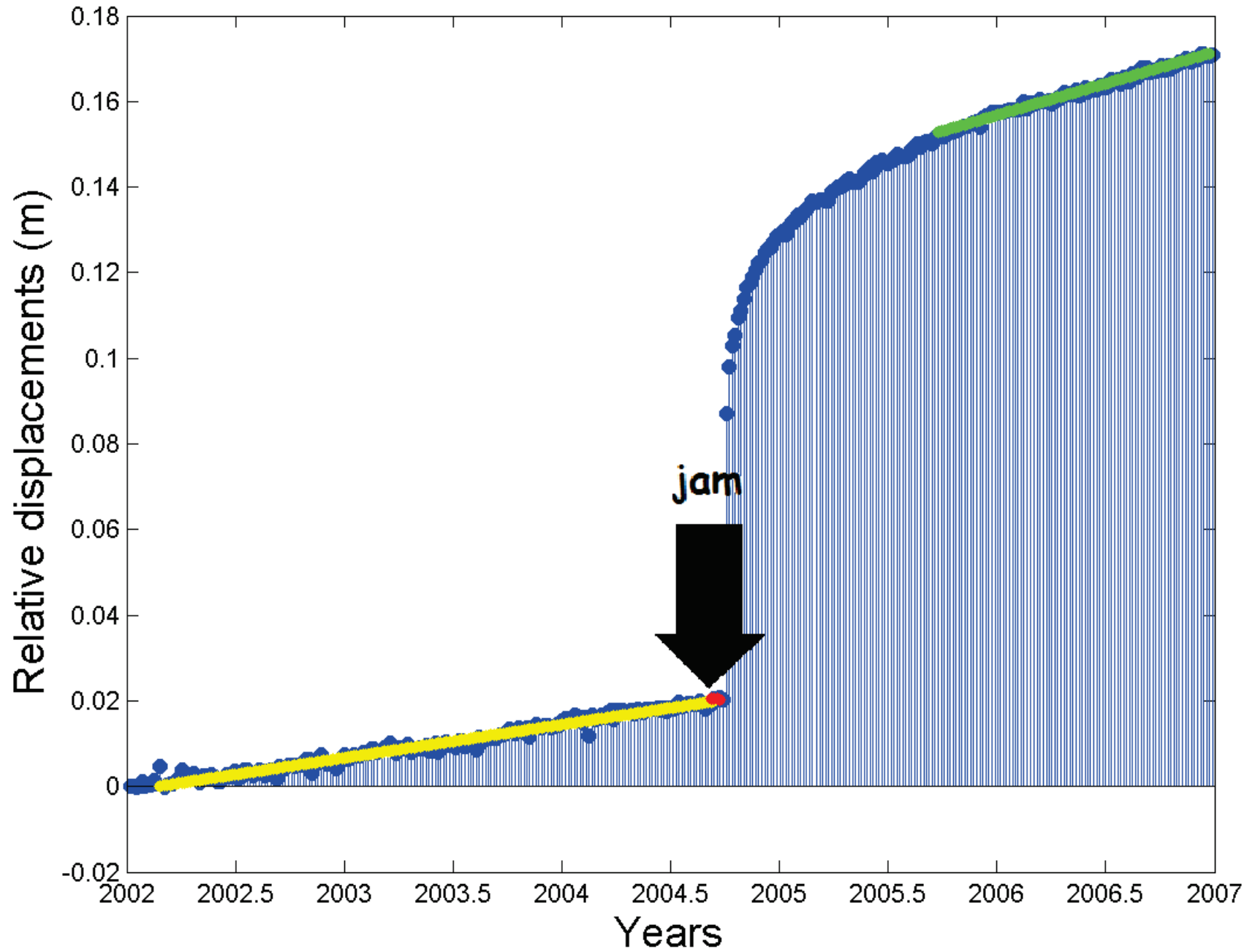


# carh-hunt

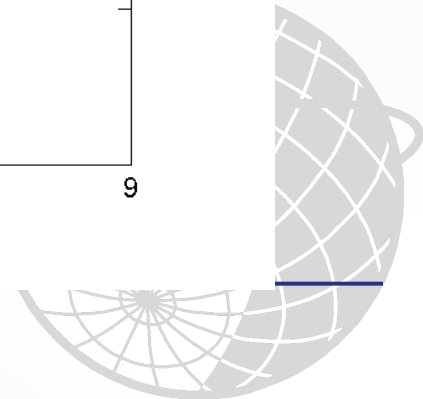
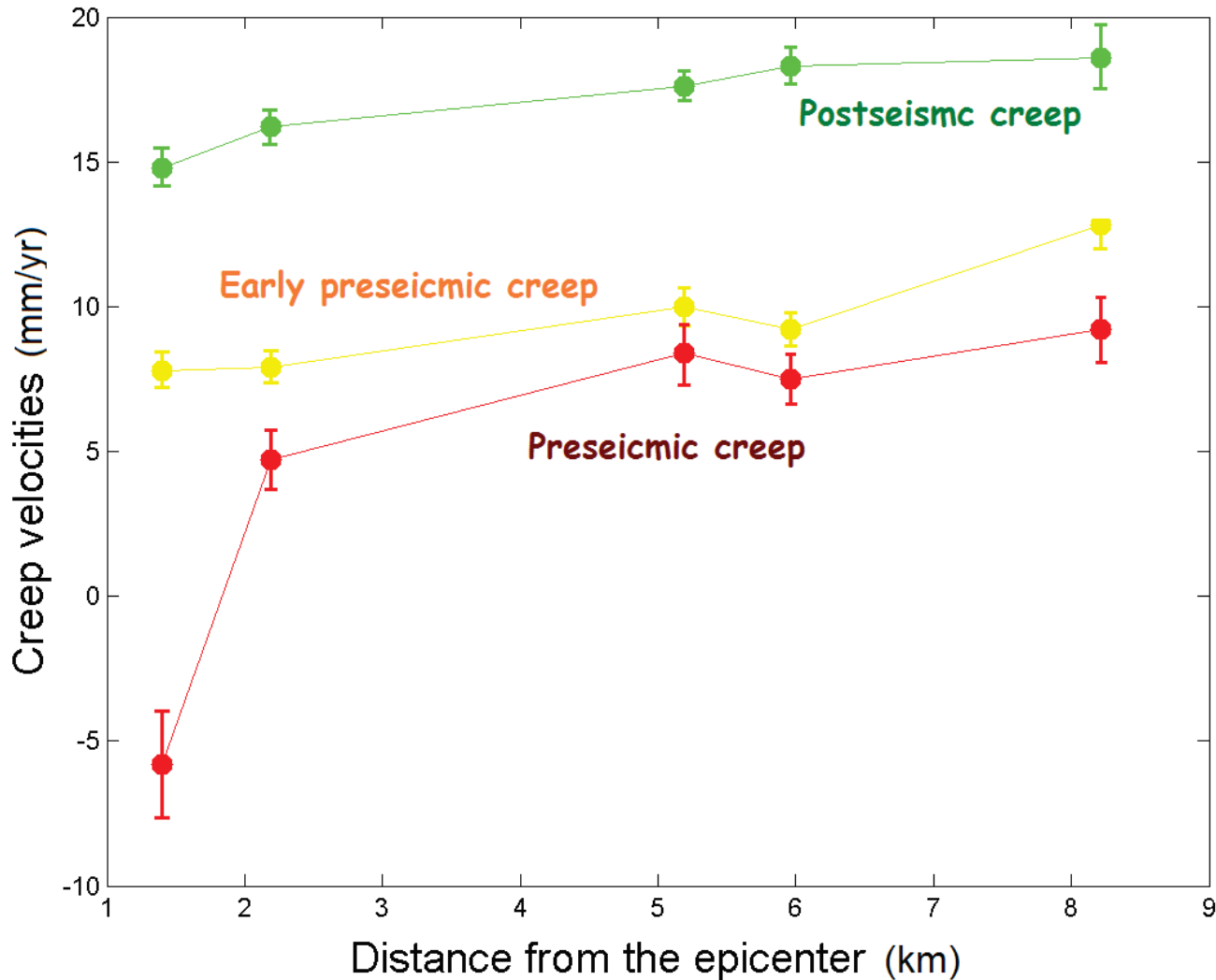




# masw-hunt

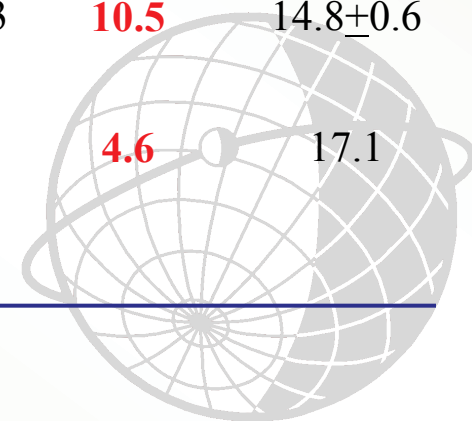


# Creep velocities decreasing approaching to the epicenter



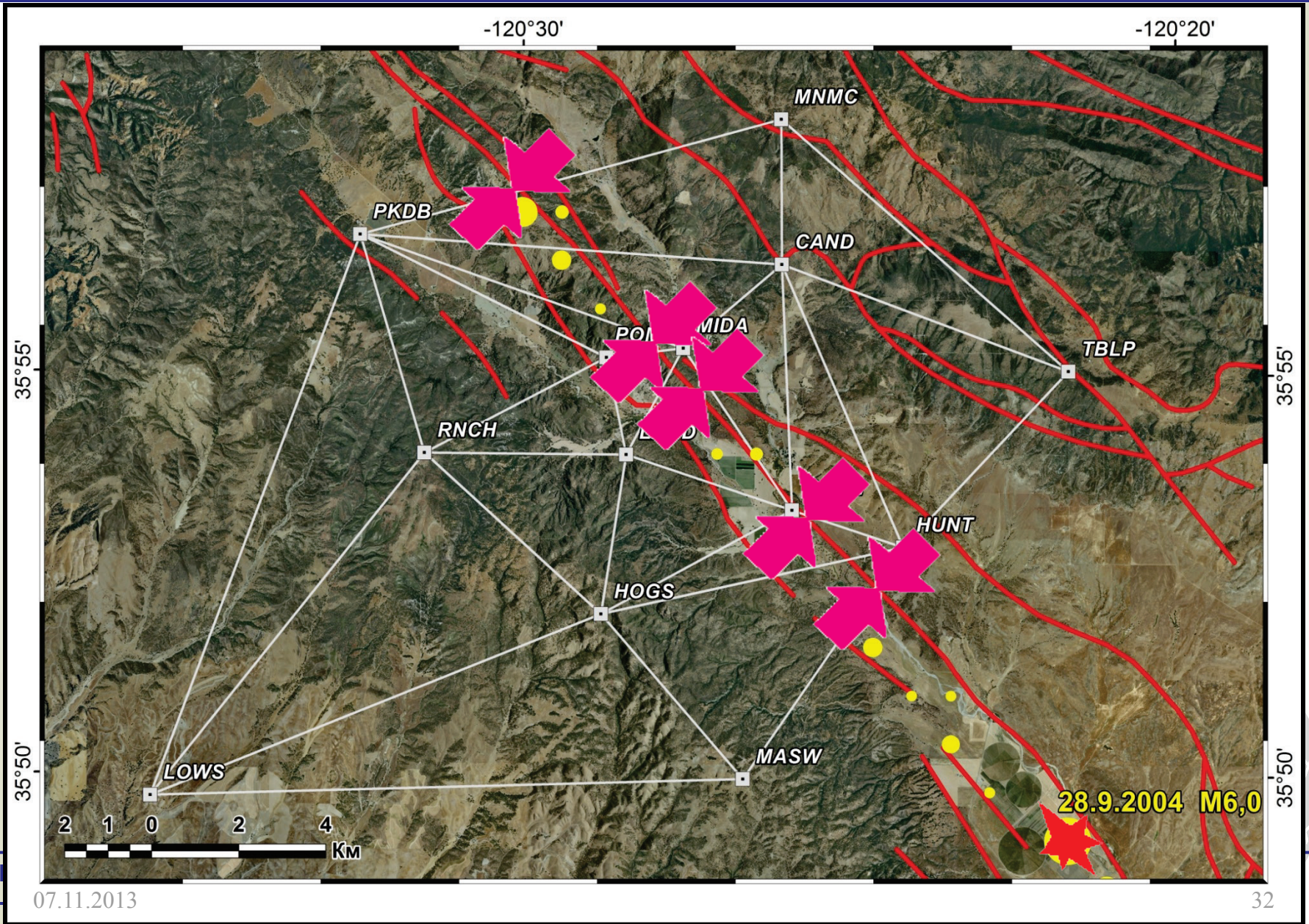
# Creep changes from a north to a south of the network

Baseline	Creep change (jam) moment	Distance to the epicenter (km)	Jam rate km/yr	Earlier interval velocity (mm/yr)	Later interval velocity (mm/yr)	$\Delta_v/\sigma_\Delta$	$t$	Post seismic interval velocity (mm/yr)
<b>pkdb-mnmc</b>	2003.9	8210		12.8 $\pm$ 0.8	9.2 $\pm$ 1.1	3.6/1.0	<b>3.6</b>	18.6 $\pm$ 1.1
<b>pomm-mida</b>	2004.2	5963	7.5	9.2 $\pm$ 0.5	7.5 $\pm$ 0.8	1.7/0.7	<b>2.4</b>	18.3 $\pm$ 0.6
<b>land-cand</b>	2004.4	5192	6.0	10.0 $\pm$ 0.6	8.4 $\pm$ 1.1	1.6/0.7	<b>2.3</b>	17.6 $\pm$ 0.5
<b>carh-hunt</b>	2004.55	2193	9.3	7.9 $\pm$ 0.5	4.7 $\pm$ 1.0	3.2/0.8	<b>4.0</b>	16.2 $\pm$ 0.6
<b>masw-hunt</b>	2004.7	1400	8.5	7.8 $\pm$ 0.6	-5.8 $\pm$ 1.8	13.6/1.3	<b>10.5</b>	14.8 $\pm$ 0.6
<b>Mean</b>			7.8	9.5	7.4*		<b>4.6</b>	17.1





# Consecutive fault locking





# Justification of elastic rebound model

- Creep velocity can be non zero just before an earthquake as it needs to be from the standard model?
- Creep velocity can be non stable months or years before an earthquake
- Creep velocity can decrease coming nearer to epicenter
- Creep velocity is about two times lower for pre- than post-seismic periods
- Elastic strain accumulation can begin not far from a previous earthquake





# Concluding remarks

- 1) Presented results demonstrate geodetic capability for modeling of earthquake preparation and realization process.
- 2) Permanent GNSS observation can be effective for searching of locked or active zones of seismic active faults.
- 3) Creep analysis during permanent observation can provide evaluation of the risk of the next earthquake occurrence.

# Thank you for attention

